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A method to quantify and account for the hygroscopic effect in stem diameter variations

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Dendrometers recording stem diameter variations (SDV) at high-resolution are useful to assess trees' water relation since water reserves are stored in the elastic tissue of the bark. These tissues typically shrink during the day as they release water when evaporative demand is high and swell during the night as they are replenished when evaporative demand is low, generating the typical SDV profile known as the diel SDV cycle. However, similar SDV cycles have been observed on dead trees due to the hygroscopic shrinking and swelling of the dead bark tissues. In order to remove this hygroscopic effect of the bark, dendrometers are applied as close as possible to the living bark tissues by removing the outer dead layer, however with questionable success. In this study, we used SDV time series from 40 point dendrometers applied on dead-bark-removed mature trees to assess and quantify the remaining hygroscopic effect on individual trees. To do so, we checked SDV behavior in the cold season and explored the relation between the diel SDV cycle and changes in relative humidity (RH). Our results showed that (a) the hygroscopic effect in SDV can be well-detected based on the amplitude of the diel SDV cycle (diel SDV_{ampl}) and the correlation between SDV and RH during both the cold and the warm season; (b) the level of the hygroscopic effect varies strongly among individuals; (c) diel SDV_{ampl} is proportional to both changes in RH and transpiration so that the hygroscopic effect on the diel SDV cycle can be quantified using a linear model where (diel SDV_{ampl}) is a function of RH changes and transpiration. These results allow the use of the model to correct the amplitude of the diel SDV cycles and suggest that this method can be applied to other ecological relevant water-related SDV variables such as tree water deficit.

KEYWORDS

hygroscopic effect, stem diameter variations, *Larix decidua*, *Picea abies*, dendrometer, diel SDV cycles

1. Introduction

Automatic dendrometers have been widely used to assess trees' water balance (Zweifel et al., 2005; Deslauriers et al., 2007; Corell et al., 2014; Mencuccini et al., 2017; Barraclough et al., 2020; Epron et al., 2021). Due to their ability to measure stem diameter variations (SDV) at micrometer resolution and at sub-hourly frequency, dendrometers not only monitor stem radial growth variations but also capture subtle water-related size variation originating from reversible shrinkage and swelling of dead and living tissue underneath the dendrometer rod (e.g., De Swaef et al., 2015).

The underlying mechanisms of those subtle variations are intimately related to trees' water use and storage and thus the trees' water status (Klepper et al., 1971). Specifically,

when canopy water loss overcomes soil water absorption, the intensifying negative water potential calls water from the stem elastic tissues (Zweifel and Häsler, 2000), which lose turgor and shrink (King G. et al., 2013). On the opposite, water refills the elastic tissues when the conditions are inverted.

This mechanism has been used to explain the daily shrinkage and swelling observed in summer, known as the diel SDV cycle (Zweifel et al., 2001; Gall et al., 2002) and to quantify trees' water deficit (TWD) as observed during extended drought periods (Zweifel et al., 2005; Oberhuber et al., 2015; Brinkmann et al., 2016; Barraclough et al., 2020). Consequently, SDV have often been used to model radial water movements in trees (Hölttä et al., 2006; Mencuccini et al., 2013, 2017; Chan et al., 2016), to assess trees' physiological drought (e.g., Martínez-Sancho et al., 2022), and to compare the impact of recent heat waves across species and space (e.g., Salomón et al., 2022).

However, these SDV-based models have been challenged by observations showing that SDV are not only related to tree-internal water use but also to the hygroscopic properties of the dead tissues of the outer phloem and bark (Lövdahl and Odén, 1992; Herzog et al., 1995; Gall et al., 2002). Even a few millimeters of dead bark in Scots pine (*Pinus sylvestris*) evoked important hygroscopic responses (Oberhuber et al., 2020) to relative humidity (RH) that differs in relation with total bark bulk density and porosity (Ilek et al., 2021). Hygroscopic responses have also been observed on the phloem of Norway spruce (*Picea abies*), probably due to the presence of dead cells in the non-conductive phloem (Gall et al., 2002). Thus, to minimize this bias, dendrometers are usually placed on the stem after removing the dead bark (Zweifel and Häsler, 2000; Turcotte et al., 2011; Oberhuber et al., 2014; Jiang et al., 2015), however with differing impact depending on: the removal success (Oberhuber et al., 2020); the species; bark, phloem and wood anatomical structure (Gall et al., 2002); and the growth rate (Zweifel et al., 2014). Moreover, as the inner structure of the bark and the proportion of living and dead bark will change over the course of a tree's life, the hygroscopic effect might vary along with age in the same individual (Krogell et al., 2012).

These considerations may pose a real challenge to the use of SDV-based models and are often not fully addressed in ecophysiological studies. The assessment of potential hygroscopic effects is especially important for big datasets in which each tree is susceptible to express different levels of hygroscopic responses. In this study we aim to provide an analytical approach to quantify and remove the impact of bark hygroscopicity on the diel SDV cycle, at the level of individual trees. The study is based on trees of both *Picea abies* and *Larix decidua* that have been equipped with automatic point dendrometers. We achieved this in two steps. First, we assessed the extent of hygroscopicity within the dataset by comparing SDV and relative humidity at different time scales and for different seasons. Hygroscopicity detection was based on the following assumptions: (1) During the cold season, when transpiration is reduced or inactive and water-related SDV are minimal, important SDV and SDV-RH correlation are signs of hygroscopicity; (2) The SDV-RH correlation coefficient is strongly associated with hygroscopicity (Oberhuber et al., 2020). Therefore a high variance for this coefficient must reflect different levels of hygroscopicity; (3) The amplitudes of the diel SDV cycles are also associated with hygroscopicity (Oberhuber et al., 2020). Therefore,

they must be correlated to the SDV-RH correlation coefficient and a high variance for this variable must also reflect different levels of hygroscopicity.

Secondly, we developed a simple model to quantify the hygroscopic effect. This model was based on the hypothesis that diel SDV amplitudes (SDV_{Ampl}) are proportional to transpiration and hygroscopicity, where the latter is in turn proportional to a change in RH. We therefore fitted a linear model for each tree, expressing diel SDV_{Ampl} as a linear combination of transpiration and diel RH amplitudes. We used the latter term as a measure of the hygroscopic effect, allowing us to apply a correction to SDV_{Ampl}.

2. Materials and methods

2.1. Study site

SDV were measured at the Lötschental tree-growth monitoring transect (Moser et al., 2010; King G. et al., 2013; King G. M. et al., 2013), which consists of 8 sites dispersed along two elevational transects of 900 meters (Figure 1) located in the Lötschental valley, Switzerland. An additional site is located just outside the valley at a lower altitude. At each site, four trees (sometimes 6) of both *P. abies* and *L. decidua* are equipped with point dendrometers. The Lötschental valley is a temperate alpine forest with a mean annual temperature of 5.7°C and a mean annual precipitation of 825 mm (King G. et al., 2013).

2.2. Data collection and processing

SDV were obtained using high-resolution point dendrometers (Radius dendrometer, DR series, Ecomatik, Germany) mounted at about 1.5 m stem height after bark removal and with a measurement frequency of 15 min from the end of 2006 to 2019 (except for trees at two sites which were instrumented in 2008 and 2013). Trees were also equipped with a Granier's Thermal Dissipation Probe (TDP) providing readings at a 15 min rate from 2012 to 2019 from April to November, with the exception of the trees at N22, N19, and N16 (Figure 1). Additionally, each site was equipped to monitor important microclimatic variables as stem temperature (with a Onset Tibit logger located in proximity to the dendrometer, 15' frequency, 2012–2019), air temperature and relative humidity (with a Hobo T-RH logger located beneath the canopy, 15', 2008–2019), as well as soil temperature and soil matrix potential (with Decagon MPS sensor located at a depth of 10 cm in the soil, 60', 2006–2019). The analyses of this study were performed using air temperature measurements. The few gaps in the time series were filled using the stem temperature measurements when available, as they consistently showed almost identical values. The SDV time series were cleaned by removing obvious outliers and abrupt changes linked to dendrometer readjustment on the stem (see Krogell et al., 2012) whereas TDP time-series were processed using the TREX R package (Peters et al., 2021) to obtain sapflux density according to the "predawn" method (see Peters et al., 2018). Before performing further analyses, all the timeseries including the environmental variables were quality checked.

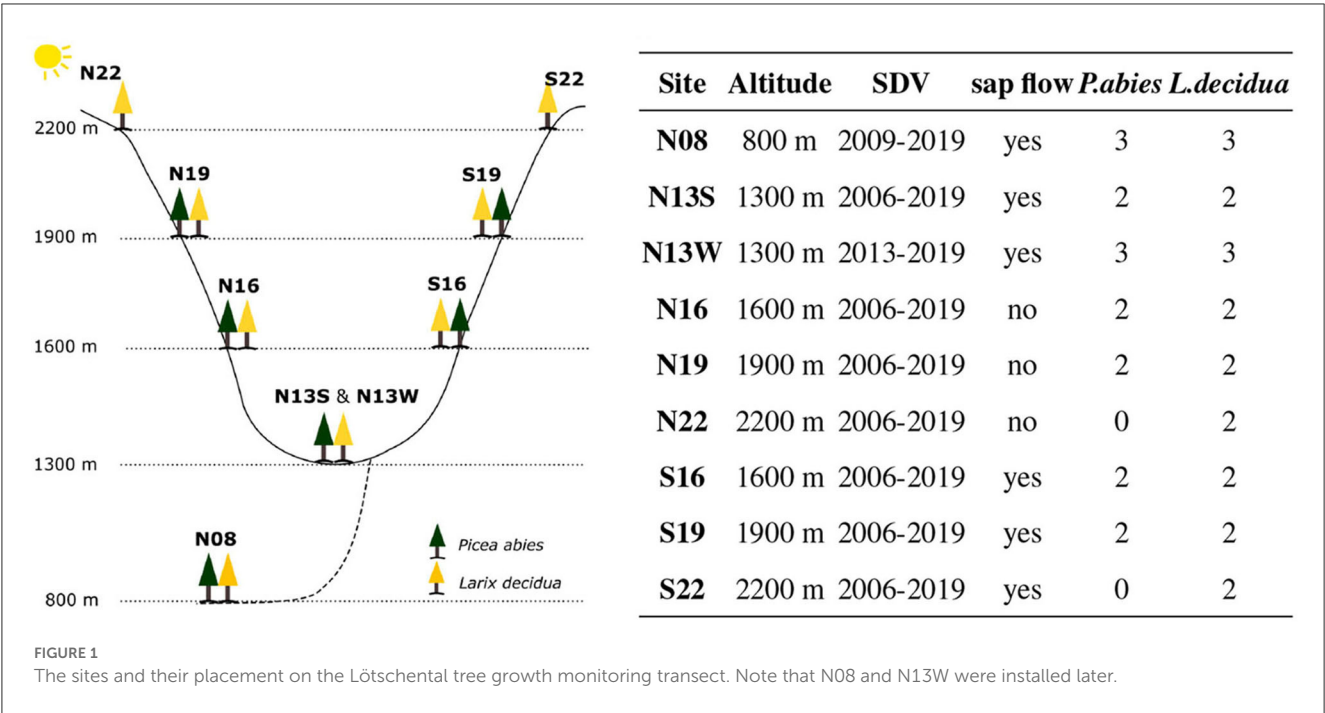


TABLE 1 The different models fitted to the SDV_{Ampl} of the individual trees.

Model name	Formula
RH	$SDV_{Ampl} = a + b \cdot RH_{Ampl}$
RH + max SFD	$SDV_{Ampl} = a + b \cdot RH_{Ampl} + c \cdot \max SFD$
RH + log(VT)	$SDV_{Ampl} = a + b \cdot RH_{Ampl} + c \cdot \log(VT)$
RH + VTa	$SDV_{Ampl} = a + b \cdot RH_{Ampl} + c \cdot VTa$
RH + mean temp.	$SDV_{Ampl} = a + b \cdot RH_{Ampl} + c \cdot \text{mean temp.}$

SFD, sapflux density; VT, variable of transpiration; VTa, modified version of VT with VT set to 0 for *L. decidua* during its inactive season.

2.3. Data analysis

2.3.1. Hygroscopicity detection in winter

Hygroscopic detection in winter was performed by analyzing the correlation between SDV and RH outside of freeze-thaw cycles. Freeze-thaw cycles are characterized by a very abrupt contraction during the freezing of the apoplastic water and decontraction upon thawing, masking any relation between SDV and RH (Figure 2). They were detected using both temperature and time thresholds. The temperature thresholds relates to the temperature where freezing and thawing are susceptible to occur. They were set at -5 and 0°C, which corresponds to the freezing and thawing temperature of apoplastic water, respectively (Zweifel and Häslér, 2000; King G. et al., 2013). The time threshold relates to the time necessary for both freezing and thawing to occur under their respective conditions. A time threshold of 12 h was used for both freezing and thawing in order to capture only the most significant freeze-thaw events.

Then, the RH and SDV curves were compared during the cold season, between November and June. We first compared the general trend of both SDV and RH during the whole season and

then selected a freeze-thaw free period to perform a correlation between RH and SDV. The correlation between RH and SDV was then calculated for the freeze-thaw free period using detrended SDV (to account for the potential interference of cellular growth). Longer term trends for both RH and SDV were derived using Fast Fourier Transform (FFT) (R package *itsmr*, Weigt, 2022), and by applying a low pass filter, cutting the frequencies above a given threshold. The thresholds were set to 0.9% for the whole cold season and 5% for the rest of the analysis. SDV detrending was performed using the R package *pracma* (Borchers, 2022) by subtracting the linear trend computed based on the least-squared fit. The SDV-RH correlation was finally determined as the maximum value of a cross correlation analysis with a maximum interval of 12 h in order to account for the lagging nature between SDV and RH.

2.3.2. Hygroscopicity detection in summer and diel amplitudes

To further assess the extent of hygroscopicity in the data set, we analyzed the relation between SDV and RH during the warm season (June, July, and August). We used both the SDV-RH correlation coefficient and the ratio of the amplitudes of the diel cycles (SDV_{Ampl} : RH_{Ampl}) as proxies of hygroscopicity (Figure 3). The SDV-RH correlation coefficient was determined as the maximum correlation coefficient of a cross correlation within a time window of 72 h, in order to account for potential lagging effects between the two time series. The diel amplitudes were extracted from SDV or RH time series as the difference between the maximum and the minimum daily values. To do that, daily peaks and valleys were identified. The highest peak (maximum) was then compared to the lowest valley (minimum). In the diel SDV cycle, the maximum occurs first (usually in the early morning) and the minimum occurs later (usually in the evening). In order to avoid measuring

amplitudes in days where the diel cycle is not properly defined, we discarded amplitudes not respecting this order.

We also compared the ratio $SDV_{Ampl} : RH_{Ampl}$ to tree diameter and annual growth to test if the quantity of newly formed and/or older tissues has an impact on the diel SDV cycle.

2.3.3. SDV amplitudes modeling

To derive the portion of SDV_{Ampl} that is under the influence of the hygroscopic effect, we fitted linear models for each tree, where SDV_{Ampl} is expressed as function of RH_{Ampl} as well as transpiration (Table 1). These models are based on the assumption that the hygroscopic effect is proportional to RH_{Ampl} .

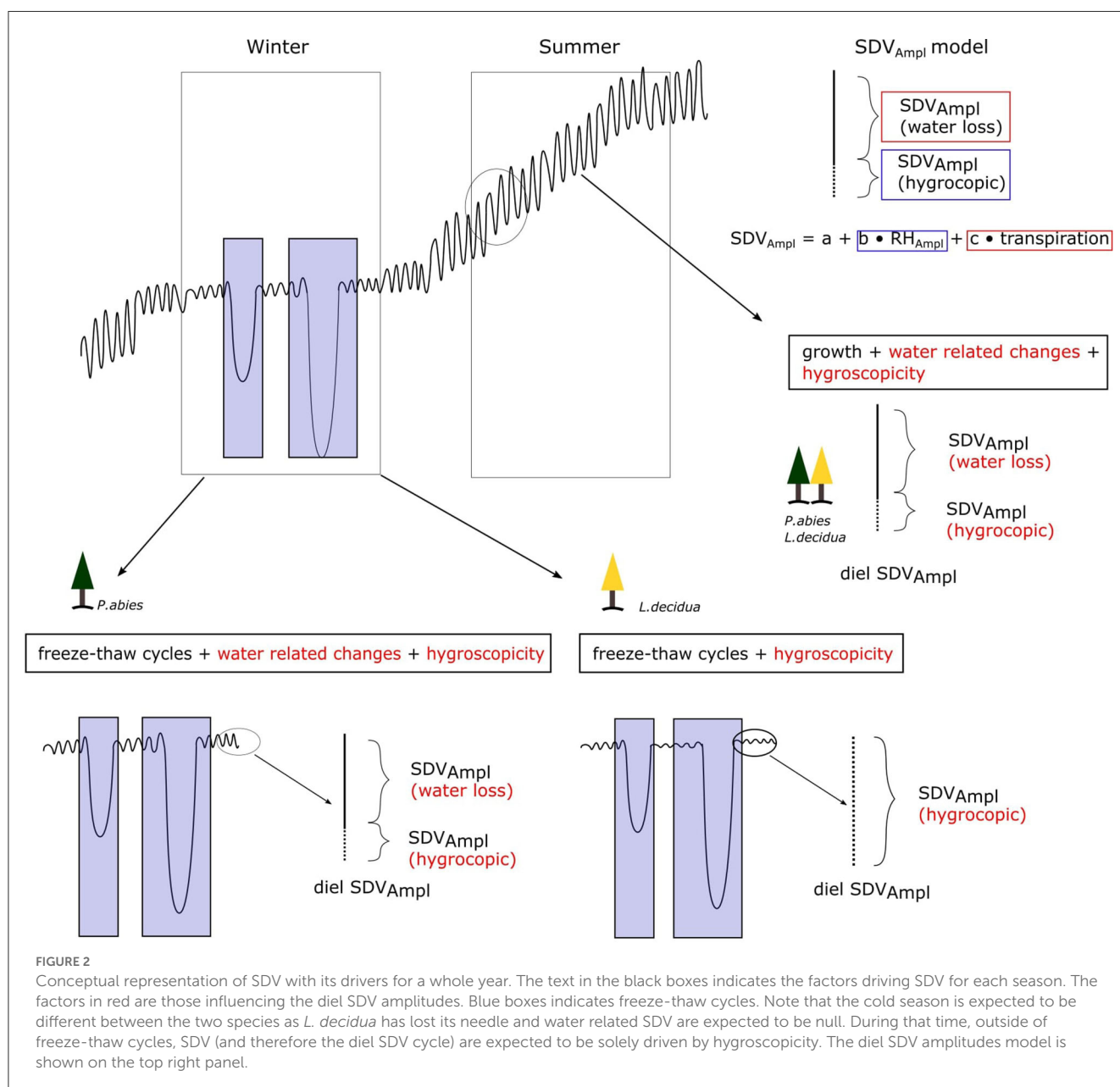
Transpiration was estimated using three different variables: diel maximum sapflux density (max SFD), the diel mean temperature,

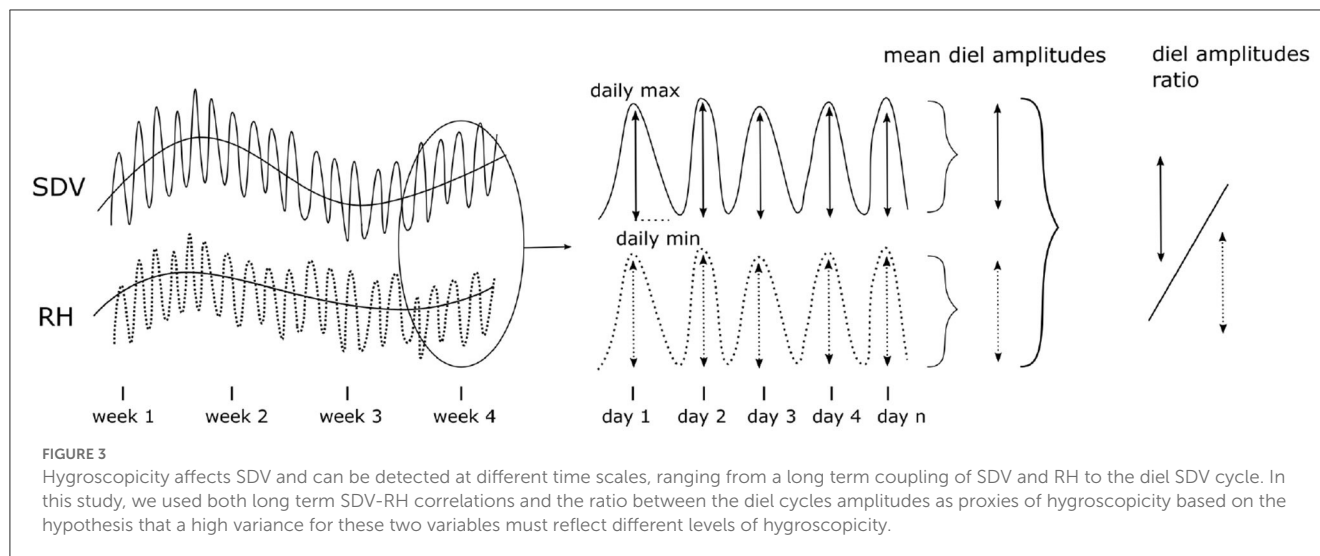
and the variable of transpiration (VT). Max SFD is the daily max value of SFD. The diel mean temperature is the mean temperature between the highest point (peak) and the lowest point (valley) of the SDV diel cycle. VT is an integrated variable including both daily solar radiation R_s and VPD and has been used as a proxy for transpiration (Kakubari and Hoskawa, 1992; Du et al., 2011). It was calculated on a daily basis using the daily mean VPD (dVPD) and the equation :

$$VT = dVPD \times R_s^{1/2}$$

Daily solar radiations were modeled using the Annandale model which is derived from the Hargreaves and Samani model and corrected for altitude (Annandale et al., 2002) :

$$R_s = R_a \cdot A_{mod}(1 + 2.7 \times 10^{-5} \cdot Z)(T_{max} - T_{min})^{0.5}$$





With R_a being the extraterrestrial radiation; T_{\min} and T_{\max} , the maximum and minimum daily air temperature; Z , the average altitude above sea level; A_{mod} , an empirical argument set to 0.263 (Suwarno and Rohana, 2017).

We also used a modified version of VT, VTa, where VT was set to 0 for *L. decidua* during its inactive period (see below), in order to account for the needle loss.

Since *L. decidua*'s transpiration is severely limited in winter due to needle loss, we determined the transpiration active and inactive season of *L. decidua* trees. We used both phenological observations *in situ* and sap flow data to set the time span of the inactive period. Specifically, it is bordered by two phenological observations: a 100% yellowing of needles at its start and the first observation of budbreak at its end. These phenological observations were adjusted with sapflow data when available to ensure that transpiration is effectively shut down between the two events.

In the study, the inactive season is also referred to *P. abies* for the sole purpose of comparing SDV between the two species during that specific time period. Therefore, it does not relate to any phenological and/or physiological change in *P. abies*.

The linear models were created using all the data points available for each tree (with a minimum of 150 data points). Both the normality and homoscedasticity of the residuals were tested for each tree. Some slight deviations were observed, but were judged as acceptable.

2.3.4. SDV correction

The linear models were finally used to apply a correction to the mean SDV_{Ampl} . The hygroscopic effect to any given change in relative humidity (ΔRH) is calculated using the hygroscopic component of the SDV modeling equation: ΔRH and the relative humidity slope (b) (Figure 2):

$$\text{Hygroscopic effect} = b \cdot \Delta RH$$

In the case of mean SDV_{Ampl} (see Figure 2):

$$\text{mean } SDV_{\text{Ampl}}(\text{hygroscopic}) = b \cdot \text{mean } RH_{\text{Ampl}}$$

Corrections were applied to the mean SDV_{Ampl} of each tree in the data set, using all the data points available for each (all years combined).

3. Results

3.1. SDV behavior in winter

In general, the relationship between stem diameter variations (SDV) and relative humidity (RH) is strongly variable among trees. This is exemplary shown for two *L. decidua* trees in Figure 4. The first tree (Figure 4A) shows a pattern of important SDV in winter along with a strong SDV-RH relation, resulting in a high correlation coefficient in short and long term (Figures 4A, B). The correlation is still important from March to mid-April, when the soil temperature at 10 cm was slightly below 0°C (indicated by the gray boxes) (Figures 4B, C). Ten days rolling correlation profile shows a high overall SDV-RH correlation with only a slight decrease from summer to winter. The second tree (Figure 4B) does not show important SDV in winter outside of the frost-thaw events, with a maximum range of about 100 μm between March and May (compared to 400 μm for the previous tree; Figures 4A, B). The correlation between SDV and RH in late winter/early spring (March to mid-April) is low (Figures 4B, C), even though soil temperature never reached sub-zero temperature. Yearly rolling correlation shows an overall smaller SDV-RH correlation, characterized by a sharper decrease from summer to winter (Figure 4C).

3.2. Diel SDV amplitudes and correlation

Mean diel SDV amplitudes (mean SDV_{Ampl}) ranged between 50 and 200 μm in the data set (Figure 5A). The SDV-RH correlation coefficients showed a clearly increasing trend with increasing mean SDV_{Ampl} (Figure 5A). Furthermore, we found a clear linear relation between the ratio mean SDV_{Ampl} : mean RH_{Ampl} and SDV-RH correlation (Figure 5B). Mean SDV_{Ampl} were not related to

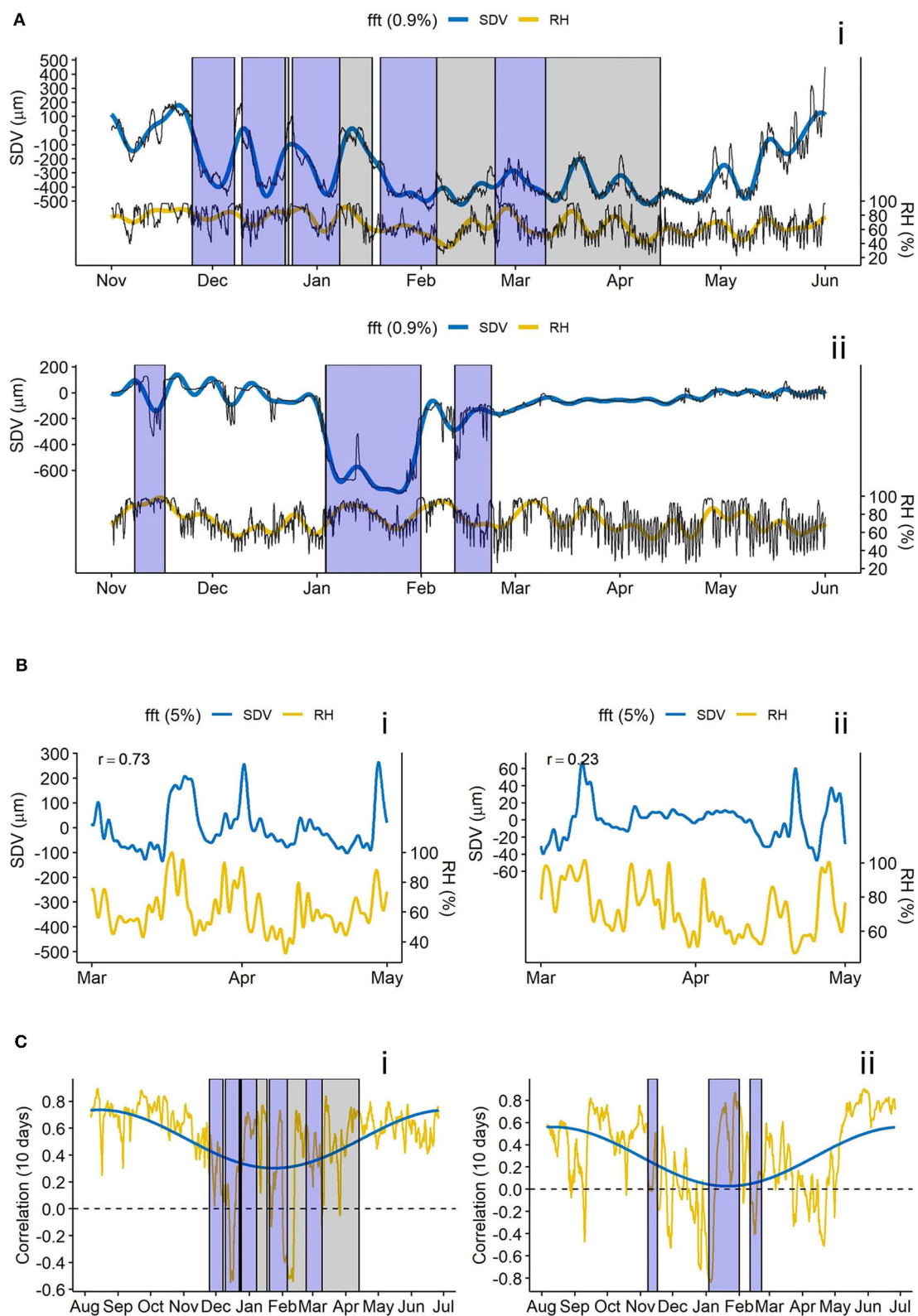


FIGURE 4

SDV - RH relations of the trees N16Ad_L2p (i = *L. decidua*, in 2010 - 2011) and N13Bd_L2p (ii = *L. decidua* in 2016-2017). **(A)** SDV and relative humidity with long term trend (fft, low pass filter of 0.9%). Blue boxes indicate freeze-thaw events, gray boxes indicate negative soil temperatures. **(B)** Detrended SDV fft and relative humidity fft (low pass filter of 5%). **(C)** 10 days rolling correlation for a year between SDV and relative humidity. Blue line shows the long term trend of the correlation (fft, low pass filter of 0.5%).

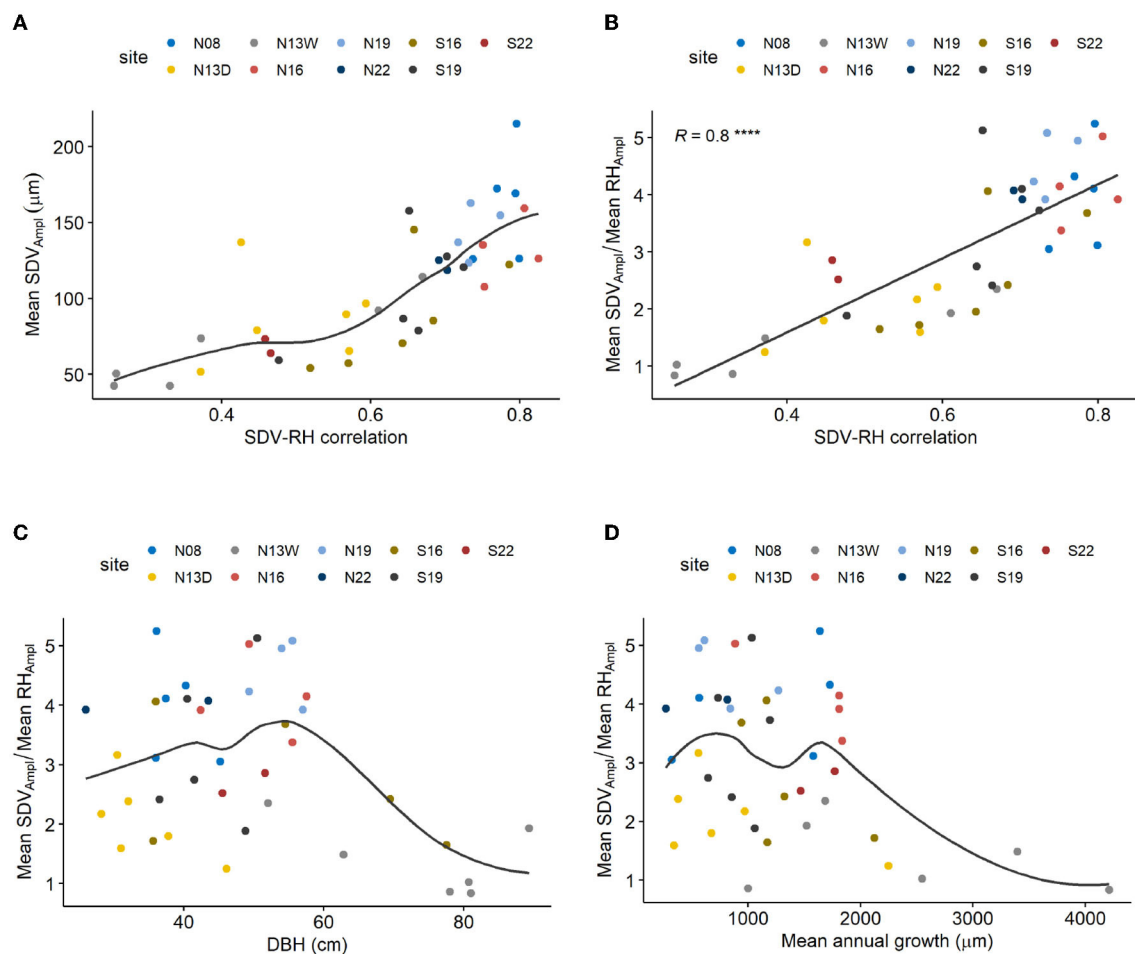


FIGURE 5

(A) Relation between mean SDV amplitude and SDV-RH correlation. Line shows Loess regression. (B) Relation between the ratio mean diel SDV_{Ampl} : diel RH_{Ampl} and SDV-RH correlation. Line shows linear regression ($****P < 0.0001$). (C) Relation between mean SDV amplitude and tree diameter (DBH). Line shows Loess regression. (D) Relation between mean SDV amplitude and mean annual growth. Line shows Loess regression. Mean diel SDV_{Ampl} - SDV-RH correlation were calculated for the summer months (June to August).

morphological variables like tree diameter (DBH) and/or annual growth (Figures 5C, D).

3.3. Diel SDV amplitude modeling and hygroscopic effect

Overall, there is a clear relation between SDV_{Ampl} and RH_{Ampl} , but with significant differences in terms of season, species and altitude. For *L. decidua*, the inactivation of transpiration has a clear effect on the correlation between SDV_{Ampl} and RH_{Ampl} (Figures 6A, B). By adding transpiration variables, the linear models for SDV_{Ampl} as function of RH_{Ampl} could be distinctly improved (Figure 7). VT and diel mean temperature resulted in highest R^2 -values. The difference between high and low altitude trees in the RH model was drastically reduced by adding VT or diel mean temperature to the model.

The residuals of the models tend to exhibit seasonal trends, with a dominance of positive values in summer and negative values in winter (Supplementary Figure 2). This tendency was especially pronounced for *L. decidua* in the RH model. Transpiration

variables, except for $\log(VT)$, were able to reduce both the temporal and the inter-specific differences. Overall, both VTa and diel mean temperature models showed the best performance. Their b coefficients are shown in Supplementary Table 1 (Supplementary material) and tend to be very similar.

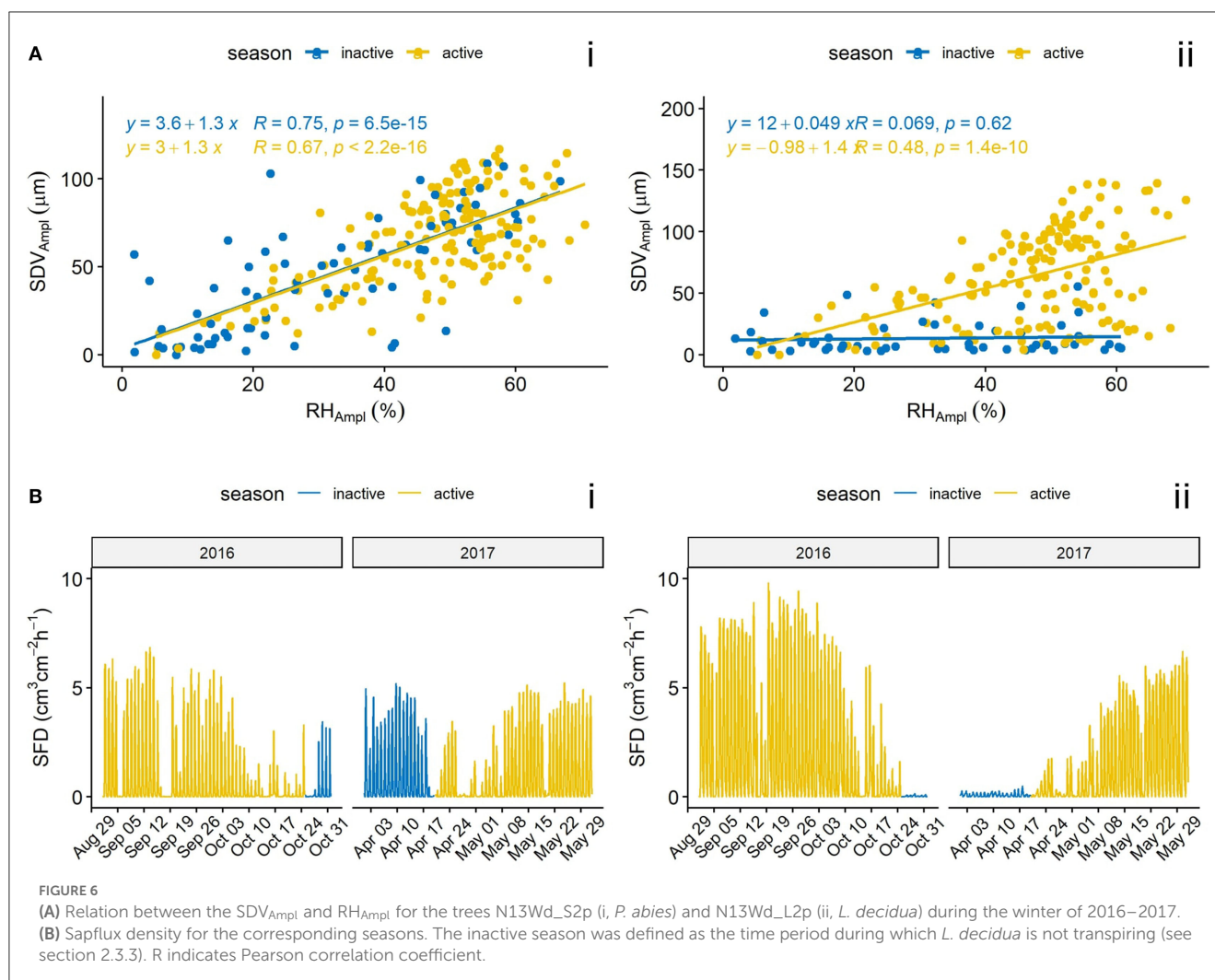
3.4. Mean diel SDV amplitudes correction

Applying the hygroscopic correction clearly reduced the mean SDV_{Ampl} and its variance among trees, similarly for both species (Figures 8A, B). The modeled hygroscopic effect shows a linear relation with the $SDV_{Ampl} - RH_{Ampl}$ ratio (Figure 9).

4. Discussion

4.1. SDV in winter and hygroscopicity

Winter proves to be an interesting season to diagnose hygroscopicity, because water related SDV are reduced (Lyu et al., 2020). In *L. decidua*'s case, at high altitudes, water related changes



should be close to zero because stomatal transpiration does not occur, and water uptake is also severely reduced upon negative soil temperature (Mayr et al., 2012; Lv et al., 2013). Despite having limited access to sapflux data in winter, they were clear signs of total inactivation of sapflux at the time of needle loss, supporting that no transpiration occurs in winter (Figure 6B). Moreover, sapflux inactivation in winter has been reported with closely related species (Urban et al., 2019).

In this context, SDV should only be shaped by the frost-thaw cycles and hygroscopic related changes. Important SDV during winter, outside the frost-thaw periods is therefore a sign of hygroscopicity. The first *L. decidua* presented in this study exhibited such a behavior (Figure 4A). Therefore, we concluded that hygroscopic effects are an important contributor of SDV for this tree.

Oppositely, the second *L. decidua* follows a SDV pattern that is closer to what is expected in winter, without hygroscopic effect (Figure 4A). To be noted that the soil temperature never reached subfreezing temperatures, indicating that water uptake might not have been impaired and therefore water related SDV might have been more important than for the previous tree. We

therefore concluded that hygroscopic effects are not an important contributor of SDV for this tree.

These two trees are located at both end of the spectrum of SDV behavior in winter at the Lötschental transect. They illustrate well the high variability of hygroscopic effects we may encounter in a big dataset. In the context of this study, many trees are exhibiting signs of hygroscopicity. We suggest therefore that assessing SDV behavior in winter might be a good way to detect hygroscopicity related SDV.

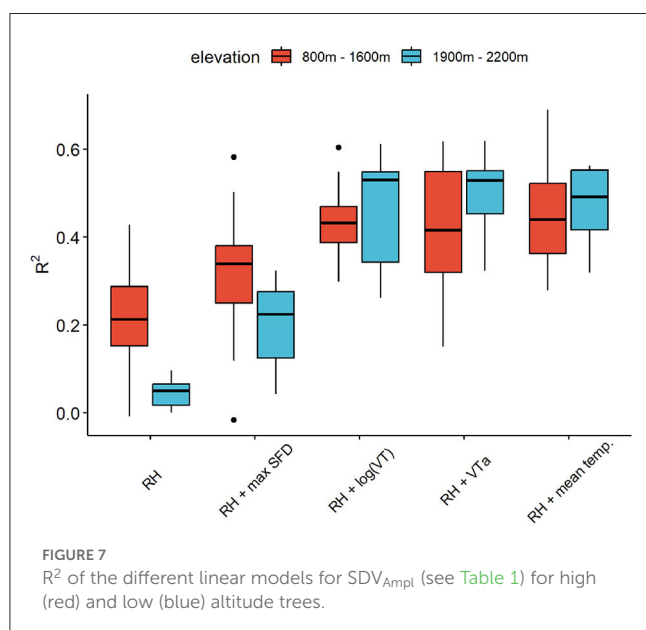
4.2. Diel SDV amplitudes and hygroscopicity

The correlation coefficient between SDV and relative humidity is a good indicator of hygroscopicity (Lövdahl and Odin, 1992; Oberhuber et al., 2020). In summer, apart from hygroscopicity, SDV express a certain degree of correlation with RH, because relative humidity is associated with the vapor pressure deficit, which is an important driver of transpiration (Zweifel et al., 2005;

Lyu et al., 2020) (Figure 6). But this correlation is expected to be lower than when induced by hygroscopicity. This is made apparent in our results by comparing yearly correlation profiles of trees (Figure 4C). Therefore high and consistent correlations between SDV and RH obtained are signs of strong hygroscopic effects. Our results indicate that the ratio $SDV_{Ampl} : RH_{Ampl}$ is a good proxy for the SDV-RH correlation and therefore hygroscopicity (Figure 5A). It is interesting to note that the ratio $SDV_{Ampl} : RH_{Ampl}$ are not related to either DBH or annual growth. If water related SDV are induced by the contraction/expansion of the elastic tissues of the living bark in response to its water content (Zweifel et al., 2000), one could expect a relation with these variables. Moreover, the variance

of the ratio $SDV_{Ampl} : RH_{Ampl}$ is similar for both species. Altogether, It suggests that most of the variance of the mean SDV_{Ampl} between individuals is influenced by the hygroscopic effect and not by water related SDV. This is on par with previous results that showed that the hygroscopic effect can be much more important than and therefore mask water related SDV (Lövdahl and Odén, 1992; Oberhuber et al., 2020). The different levels of the hygroscopic effect are, in turn, probably related to differences in the extent of dead bark removal, prior to dendrometer installation.

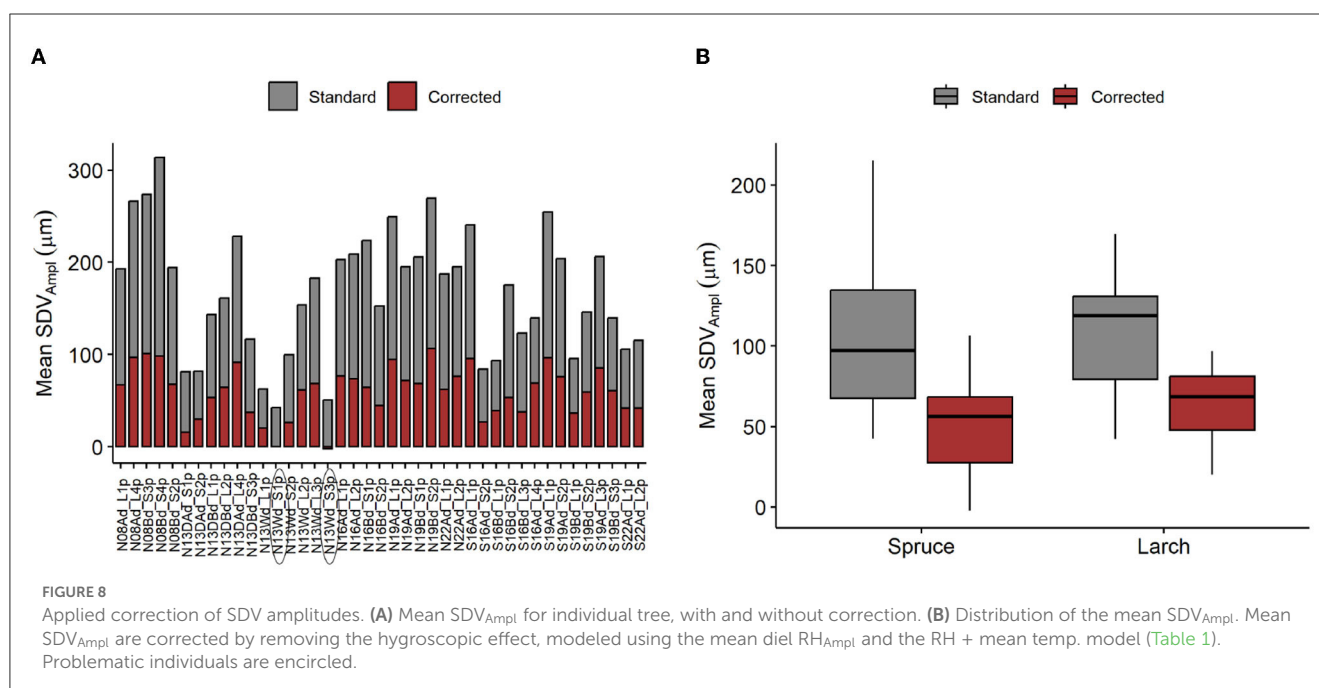
We suggest that comparing SDV_{Ampl} is valuable diagnostic tool in regards to hygroscopicity. Very high variations in SDV_{Ampl} among individuals are likely to be caused by the hygroscopic effect and should be addressed in studies using large data sets.

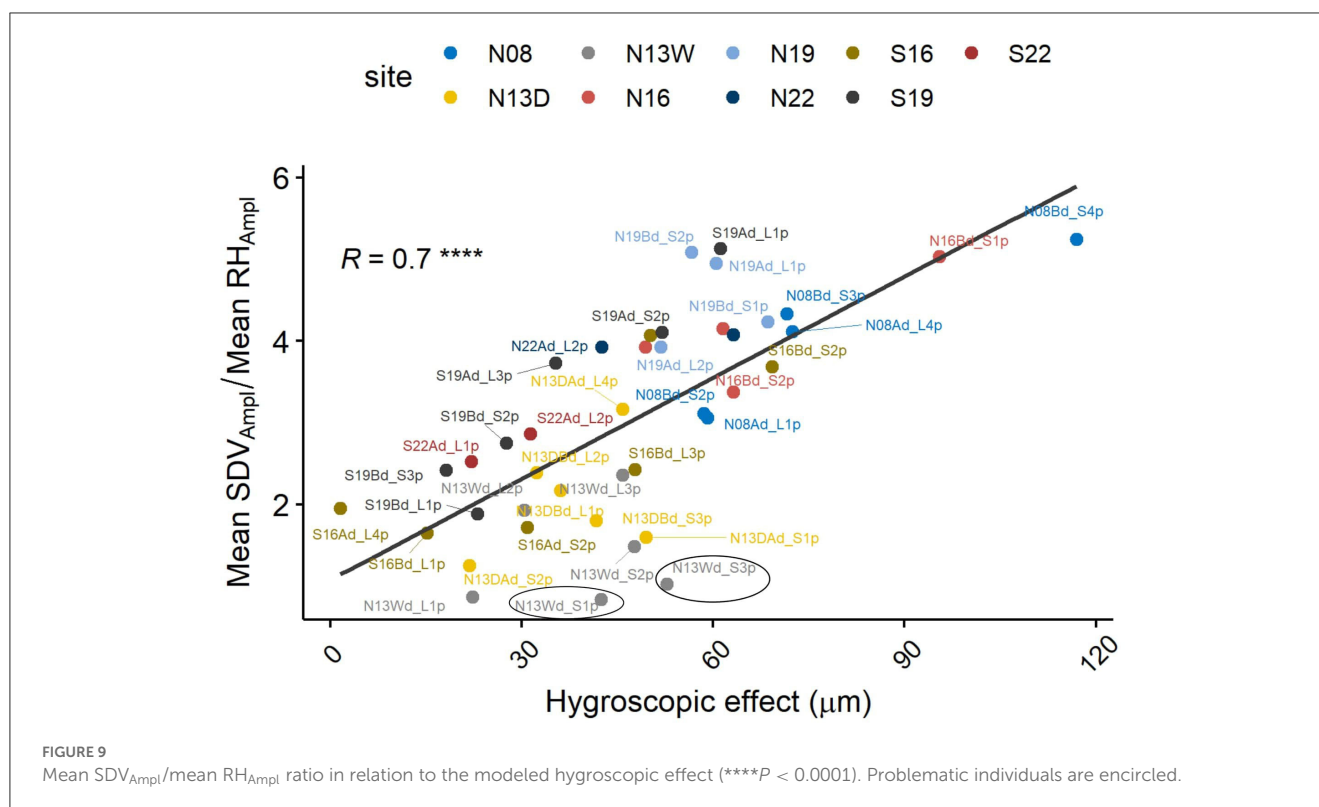


4.3. Relative humidity and transpiration influences on diel SDV

The comparison between SDV_{Ampl} and RH_{Ampl} indicates that transpiration is an important contributor of SDV. Indeed, under the hypothesis that the diel SDV cycle is shaped by relative humidity changes only (Gall et al., 2002), we would not observe the seasonal pattern for *L. decidua* that is concomitant with the activation/inactivation of transpiration (Figure 6A). This result is in agreement with the phloem water capacitance model, stating that SDV are driven by the dehydration of the living bark tissues to cover the evaporative demand of the tree (Zweifel et al., 2001).

It indicates that both relative humidity and transpiration are contributors of SDV in our dataset. As the transpiration variables and RH_{Ampl} are independent (Supplementary Figure 1C), we suggest that the relation between RH_{Ampl} and SDV_{Ampl} at the point of no transpiration is equivalent to the hygroscopic effect. Following this hypothesis, the linear regression between diel SDV and diel RH amplitudes for *L. decidua* during the inactive season





is a close approximation of the hygroscopic effect because no transpiration is detected during that time and therefore no water related SDV are occurring.

4.4. Plausability of the model

It is evident that a simple linear relationship between SDV_{Ampl} and RH_{Ampl} is not a sufficiently good estimation of the hygroscopic effect, as SDV_{Ampl} are also driven by transpiration. This is also apparent in the seasonal distribution of the residuals of the linear regression between SDV_{Ampl} and RH_{Ampl} (Supplementary Figure 2). This model is underestimating the summer amplitudes, when transpiration is high, and overestimating the winter amplitudes when transpiration is low. This effect is noticeably stronger for *L. decidua*, because this species expresses a greater seasonal difference in transpiration levels. During the inactive season of *L. decidua* the situation is different because no transpiration is detected, and water related SDV are therefore expected to be null. In that time period, the relation between SDV_{Ampl} and RH_{Ampl} must account for the hygroscopic effect only. Following this idea, the hygroscopic effect can be quantified for *L. decidua* during its inactive season. However, the prevalence of freeze-thaw cycles and the lack of data points makes it difficult in practice, resulting in very poor R^2 . Moreover, this approach would not work for *P. abies* since for this evergreen species transpiration occurs throughout the year. The purpose of adding a transpiration variable to the model was to ensure that the coefficient $b RH_{Ampl}$ would be representative of the hygroscopic effect. The whole plausibility of the model relies on the fact that

the underlying effect of transpiration on SDV_{Ampl} is successfully taken into account in the model. There are indications that this is the case. Firstly, the seasonal distribution of the residuals of the model and the differences among species are drastically reduced (Supplementary Figure 2). Secondly, the difference in R^2 between low and high altitudes sites has been drastically reduced or removed by adding transpiration to the models (Figure 7). The difference in R^2 is likely related to the fact that the range of transpiration values is larger at higher altitudes (with longer periods of total absence of transpiration). The added variance is not explained by the RH only model. Therefore, reducing these differences is an indication that the transpiration has been successfully taken into account. Thirdly, the corrected mean SDV_{Ampl} shows a noticeable lower variance than the non-corrected version (Figure 8B). We postulated that hygroscopic effects are an important contributor to summer oscillation variance between trees. The decrease in variance points therefore to a reduction of the hygroscopic effect. Together, these are good indications that the model is plausible and can be used for both species to derive the hygroscopic effect.

4.5. Model performance and intended usage

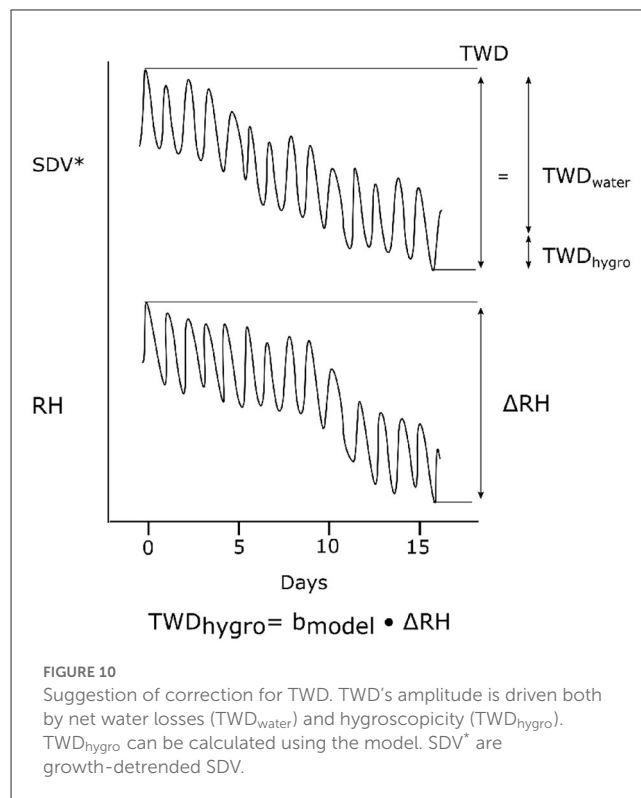
Of all the linear models, both VTa and mean temp. seem to perform better than max SFD (Figure 7). This result is surprising as SFD is a more direct measure of transpiration. The transpiration data set, however, presented noticeable imprecisions, but no cleaning was performed in this study. Both, the VTa and the mean. temp based model presented very similar coefficients for

the intercept and the relative humidity slope for many trees, but we chose to keep the mean temp. model based on slightly better performance.

In this model, the seasonal differences of the distribution of the residuals are reduced but not completely evened out, meaning that some temporal correlation remains (Supplementary Figure 2). Residuals in winter tend not only to be on the negative side but also to express less variation, causing heteroscedasticity, noticeable between the lower values of the winter oscillations and the typical higher summer oscillations residuals. In this context, heteroscedasticity means that summer oscillations are given more weight in the model. We consider that this does not represent an important issue as the diel SDV amplitudes are more robustly coupled with the diel RH amplitudes in summer. Depending on the relative number of winter data points included in the model, there can also be a slight deviation to the normality of the residuals, especially at the highest and lowest values.

This seasonal imbalance may be due to the effect of several phenomenons. Firstly, it is possible variation in water availability are affecting diel SDV_{Ampl}. In summer, especially during droughts, water uptake would be reduced, inducing higher SDV_{Ampl} for the same amount of transpiration. Secondly, carbon related processes such as leaf photosynthetic activity can affect the amplitude of the phloem water capacitance response (Steppe et al., 2015). Thirdly, the hygroscopic response is slightly temperature dependent (Lövdahl and Odin, 1992). The impact of relative humidity changes on SDV in winter might therefore be slightly smaller compared to summer. All these effects combined are negatively impacting the usability of the model. A few notable imprecisions are present. Firstly the intercept of the linear model should be theoretically around 0, but many trees deviate from that (Supplementary Figure 1 and Supplementary Table 1). Secondly, the model seems to be more limited for a few trees. For instance, the corrected mean SDV_{Ampl} of some individuals are close to zero, or even slightly negative (encircled individuals, Figure 8A). This would make sense if the trees were dead and water related changes would be null. In these cases, it is apparent that the correction is over-evaluated. However, the clear linear relationship between the hygroscopic correction and the ratio SDV_{Ampl} : RH_{ampl} is an indication that the correction applied is proportional to the level of hygroscopicity (Figure 9). It indicates that the modeled correction is generally pointing to the right direction despite having individual deviations. If that is the case, one might want to use a correction based on the linear regression expressed in Figure 8B. To establish this approach, we would need more data, preferably from other sites, to test whether there is indeed a universal linear relation between the ratio SDV_{Ampl} : RH_{ampl} and hygroscopicity. Experiments could also be done to test this hypothesis. One approach would be to control the amount of bark removal prior to dendrometers installation and test how the dead bark thickness affects the ratio SDV_{Ampl} : RH_{ampl}.

Despite the potential issues and imperfections of the models, this approach is, to our knowledge, the first to tackle the issue of hygroscopicity. We suggest that it can be used to quantify the hygroscopic effect of notable water related SDV variables like TWD. TWD is calculated from the contraction of the stem in summer in response to water deficit (Zweifel et al.,



2005). It is defined as the amplitude of the SDV contraction after removing the cellular growth. However SDV decrease over long period in summer are often concomitant with decrease in relative humidity due to the more arid conditions provoking the drought. Therefore it is not clear whether the contraction of the stem is linked with water depletion or the hygroscopic effect, if the latter is indeed present in the data. Our models could be used to derive the fraction of the contraction under the influence of the hygroscopic effect, by taking the RH_{Ampl} at the beginning and the end of the SDV decrease on which TWD is derived (Figure 10). Such a correction could help to standardize TWD and allow for a more accurate comparison between individuals.

5. Conclusion

The hygroscopic response of the bark can alter the interpretation of SDV in terms of water related changes and studies often consider it to be null with the practice of dead bark removal. In our study, the trees exhibited different levels of hygroscopic response. In winter, except for freeze-thaw events, SDV are mainly shaped by the hygroscopic response of the bark, as water-related changes are minimal. The amplitude of the diel SDV cycle is influenced by transpiration, in agreement with the phloem capacitance hypothesis, and also by relative humidity. The contribution of relative humidity on SDV corresponds to the hygroscopic effect and can be quantified by a simple linear model where diel SDV amplitudes is the response variable and both RH amplitudes and transpiration are the predictors. This kind of model can then be used to calculate the hygroscopic response for

a given change in relative humidity. We suggest this model for correcting SDV derived variables used to assess the water balance of a tree.

Data availability statement

The raw data supporting the conclusions of this article will be made available by the authors, without undue reservation.

Author contributions

FD: data analysis and modeling, conceptualization, data visualization, and writing—original draft. PF: conceptualization, methodology, data provision, and writing—review and editing. HL: conceptualization, methodology, and writing—review and editing. CM: supervision, conceptualization, methodology, and writing—review and editing. All authors contributed to the article and approved the submitted version.

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Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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Supplementary material

The Supplementary Material for this article can be found online at: <https://www.frontiersin.org/articles/10.3389/ffgc.2023.1167542/full#supplementary-material>

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